Also, we can use the Youngs-La Place equation to related pore radius to soil pressure head, or:

$$\Delta P = \rho_w gh = \frac{2\gamma \cos(\alpha)}{R}$$
 Assume $\alpha = 0$

Or, for each pore with radius $R_{j} = \frac{2\gamma}{\rho_{w}gh_{j}}$

$$R_{J}^{2} = \frac{4\gamma^{2}}{(\rho_{w}gh_{J})^{2}}$$
, or

$$K_{sat} = \frac{\rho_{w} g \Delta \theta_{J}}{8 \eta \tau^{2}} \frac{4 \gamma^{2}}{(\rho_{w} g)^{2}} \sum_{J=1}^{M} \frac{1}{h_{J}^{2}}$$

$$K_{sat} = \frac{\Delta \theta_{J} \gamma^{2}}{2\eta \tau^{2} \rho_{W} g} \sum_{J=1}^{M} \frac{1}{h_{J}^{2}}$$

$$K_{sat} = Constant \sum_{J=1}^{M} \frac{1}{h_J^2}$$

Where all values for Constant are known, except τ !!!!!!!!

Possibly, $\tau = \tau(\theta)$, as $\tau = L_c/L$ HOW ????

$$K_{sat} = Constant \sum_{J=1}^{M} \frac{1}{h_{J}^{2}}$$

where K_{sat} is controlled by total soil pore space (all pores are filled with water), hence for all pore classes J =1 (largest pores) to M (smallest pores).

Hence,

$$K(\theta_{s} - i \Delta \theta_{j}) = Constant \sum_{J=i+1}^{M} \frac{1}{h_{j}^{2}}$$

i = 0, ..., (M-1)

with i = 0 representing saturated soil, and i = 1 removing soil pores (capillaries) that drain at $h = h_1$ (corresponding with r_l , Youngs equation). Therefore, this corresponds with J = 2, M.



Exercise will ask you (a) to program the solution of unsaturated conductivity data for selected soils (for which soil water retention curves are given), and (2) to determine τ –value by fitting K-solution to known K-data.

Both soil water retention and unsaturated hydraulic conductivity values are determined by van Genuchten-Mualem hydraulic functions:



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In class we derived an expression to predict the unsaturated hydraulic conductivity from Poiseuille's law.

$$K_{s} = \frac{\Delta \theta_{J} \gamma^{2}}{2 \eta \tau^{2} \rho_{w} g} \sum_{J=1}^{M} \frac{1}{h_{J}^{2}} = Constant_{J} \sum_{J=1}^{M} \frac{1}{h_{J}^{2}}$$
(1)

which leads for the unsaturated K to:

$$K(\theta) = K(\theta_s - i\Delta\theta_J) = Constant_J \sum_{J=i+1}^{M} \frac{1}{h_J^2} \quad i = 0, \dots, (M-1)$$
(2)

with i = 0, indicating the contribution to K of all pore classes (K_s), and i=M-1 (J=M) corresponding with K at the lowest water content (only the smallest size pore class contributes).

1. Use the above expressions to predict $K(S_e)$ for the sandy and loamy soils, for which measured retention and unsaturated hydraulic conductivity data were fitted to the van Genuchten-Mualem parameters (use the van Genuchten, vG, retention model to obtain θ for a given h).

Use $\Delta \theta_J = 0.02$. List values for all parameters embedded in *Constant*, and their units. Compute τ by matching the predicted K at saturation to the listed K_s value in the table below. Thus, assume that τ is independent of water content, and hence is a constant parameter. So, for this exercise you need to use the vG retention model and Ks value from the table.

| Parameter | Sand | Loam |
|------------------------------|-------|-------|
| θ_{s} | 0.43 | 0.43 |
| $\Theta_{ m r}$ | 0.045 | 0.078 |
| α (cm ⁻¹) | 0.145 | 0.036 |
| n | 2.68 | 1.56 |
| $K_s (cm h^{-1})$ | 29.7 | 1.04 |
| m=1-(1/n) | | |
| | | |

Comment [m1]: Jan, If we use $\Delta \theta$ constant and equal to 0.02, I think here we should say "to obtain h for a given θ ". Vadose Zone Hydrology – HYD210 Due: 04/10/2015 (Before class starts) Æ

van Genuchten retention model:

$$\frac{\theta - \theta_{\rm r}}{\theta_{\rm s} - \theta_{\rm r}} = S_{\rm e} = (1 + (\alpha \mid h \mid)^{\rm n})^{-\rm m}$$
(3)

2. In addition to using the calibrated τ -value at saturation only, approximate the tortuosity dependency on water content, by matching measured with predicted K-data (from the van Genuchten-Mualem K-model) at each of the selected water content values of question 1. Hence, the parameter *Constant* is now variable, and is a function of θ . Fit the newly predicted to the following exponential expression: $\tau = a S_e^{b}$

van Genuchten-Mualem K-model:

$$K(\theta) = K_s S_e^{0.5} [1 - (1 - S_e^{1/m})^m]^2$$
(4)