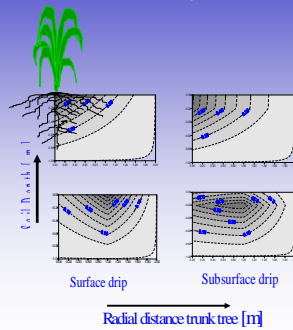


PARAMETER OPTIMIZATION USING INVERSE MODELING (soil hydraulic properties and root water uptake)



WHAT IS INVERSE METHOD ???

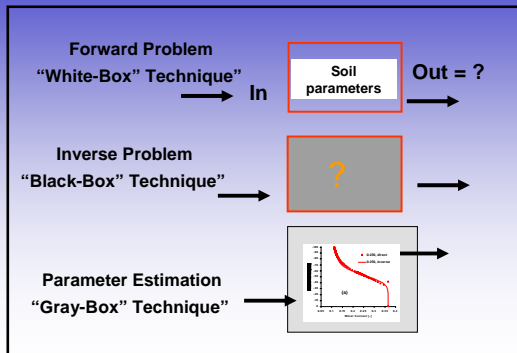
Solution of an inverse problem entails determining unknown causes, based on observation of their effects

This is in contrast to the corresponding direct problem, whose solution involves finding effects based on the complete description of their causes.

EXAMPLES:

- o Computerized Tomography (CT)
- o Boundary inverse problem and backward problem (to find initial conditions)
- o Parameter Estimation (pde parameters)

PARAMETER ESTIMATION by inverse modeling



For example, fit parameters to Convection-Dispersion Equation (Parameter estimation)

$$\text{Problem: } R \frac{\partial c}{\partial t} = D \frac{\partial^2 c}{\partial x^2} - v \frac{\partial c}{\partial x}$$

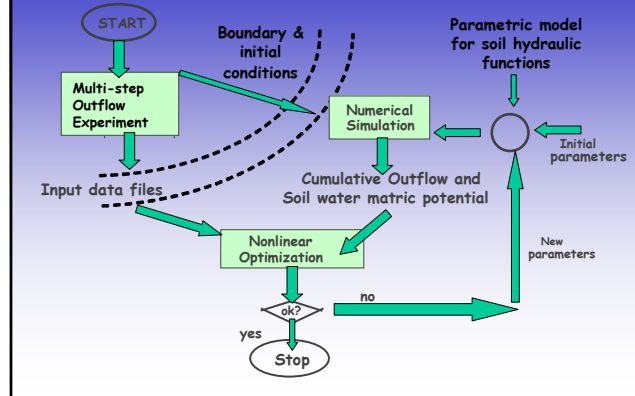
$$\text{Solution: } \frac{c}{c_0} = \frac{1}{2} \operatorname{erfc} \left[\frac{Rx - vt}{\sqrt{4DRt}} \right] + \frac{1}{2} e^{\frac{vx}{D}} \operatorname{erfc} \left[\frac{Rx + vt}{\sqrt{4DRt}} \right]$$

Difference: Solution is explicitly known, whereas in inverse modeling, solution can be only obtained by numerical modeling

WHY NEED FOR MEASUREMENT OF SOIL HYDRAULIC PARAMETERS ?

- ❑ As input to water flow and contaminant transport models;
- ❑ To characterize soil physical characteristics, including their spatial and temporal variability;
- ❑ To correlate with other, more easily to measure soil physical properties, e.g. texture.

Flowchart of Inverse Modeling (IM)



Parameter estimation of soil hydraulic properties, (Methods of Soil Analysis, 2002)

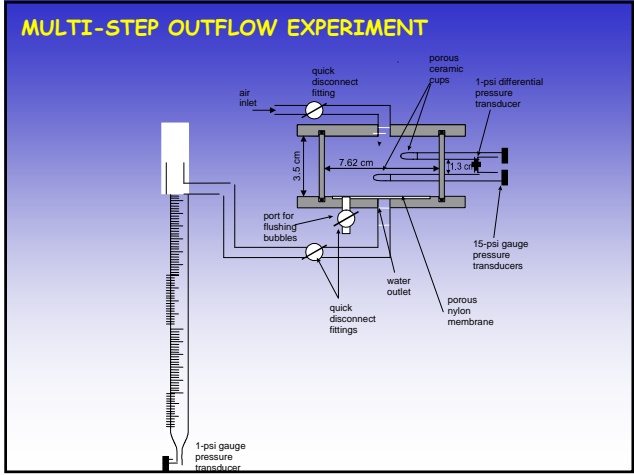


$$S_e = \left[\frac{1}{1 + (\alpha h)^n} \right]^m$$


$$K_r = S_e^l \left[1 - (1 - S_e^{1/m})^m \right]^2$$

Miniature tensiometer for multi-step experiments

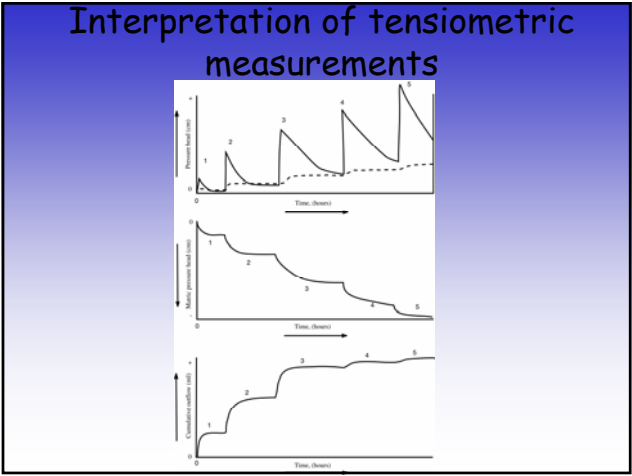
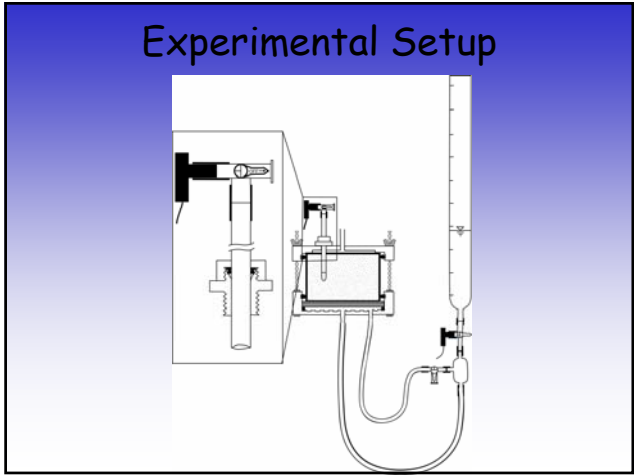




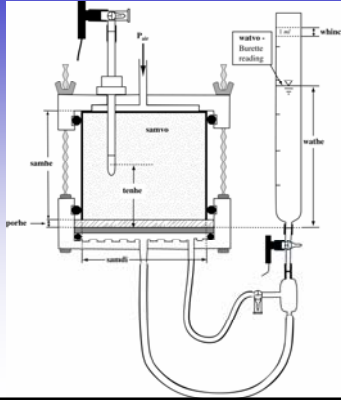
Experiment:



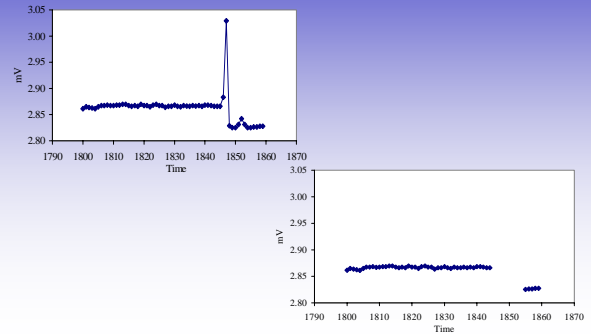
- Multi-step outflow, with tensiometric measurements inside soil core;
- Apply a sequence of air pressure steps to initially near-saturated soil core;
- Monitor cumulative drainage volume and tensiometer pressure with pressure transducers;
- Measure boundary and initial conditions



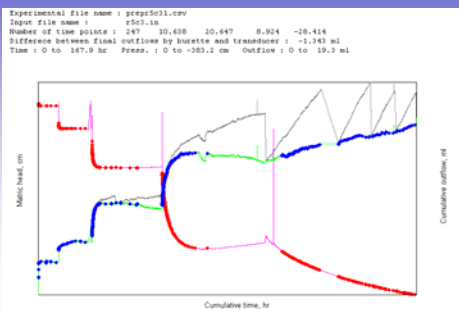
Data prep variables



Correction to outflow transducer as caused by flushing

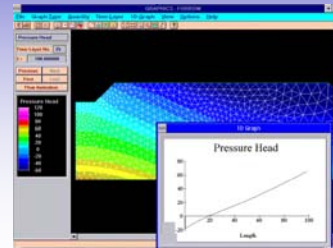


Correction with dataprep

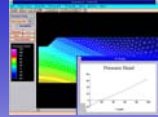


SIMULATION MODEL:

HYDRUS-2D for Windows
 Simulating Two-Dimensional Water Flow,
 Heat and Solute
 Movement in Variably Saturated Media



MODELING:

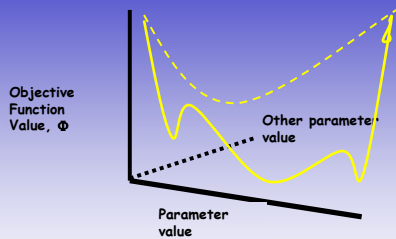


- Use van Genuchten–Mualem or Kosugi lognormal pore size distribution model to describe soil water retention and unsaturated hydraulic conductivity functions;
- Assume initial parameter values for these functions;
- Solve transient one-dimensional water flow model with known initial and boundary conditions to solve for cumulative drainage and soil water matric potential

PARAMETER OPTIMIZATION

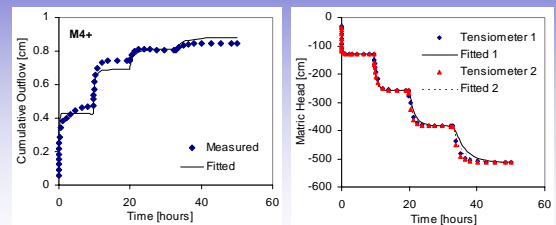
- Compare measured with simulated flow variables;
- Find optimum parameter set, so as to minimize differences between measured and simulated flow variables: I.e. soil water matric potential and cumulative drainage from soil sample
- Use Levenberg–Marquardt method, Simplex method or Genetic Algorithm

Optimization - Minimize Objective Function



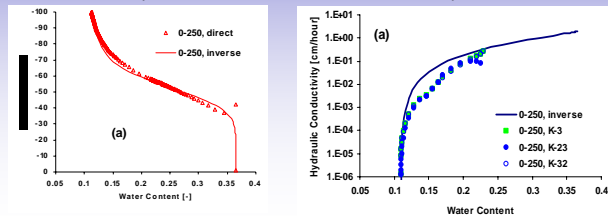
$$\Phi(\mathbf{b}) = W_Q \sum_{i=1}^N \{\omega_i [Q_{exp}(t_i) - Q_{sim}(t_i, \mathbf{b})]\}^2 + W_{h_m} \sum_{j=1}^M \{\omega_j [h_{m,exp}(t_j) - h_{m,sim}(t_j, \mathbf{b})]\}^2$$

COMPARISON OF MEASURED WITH SIMULATED OUTFLOW AND MATRIC POTENTIAL



Multi-step outflow method to indirectly estimate soil water retention and unsaturated hydraulic conductivity functions

Lincoln sand
(Wildenschild et al., 2001)



Example of multi-dimensional root water uptake
In Press: Vrugt et al., SSSAJ and Water Resour Res.

Solve unsaturated water flow equation for multi-dimensional soil domain, with root water uptake term (S)

$$\frac{\partial \theta}{\partial t} = \nabla \cdot [\mathbf{K} \nabla (h - z)] - S(x, y, z, t)$$

MULTI-DIMENSIONAL ROOT WATER UPTAKE MODEL

$$\beta(x, y, z) = \left(1 - \frac{x}{X_m}\right) \left(1 - \frac{y}{Y_m}\right) \left(1 - \frac{z}{Z_m}\right) e^{-\left(\frac{D_x}{X_m} |x^* - s| + \frac{D_y}{Y_m} |y^* - s| + \frac{D_z}{Z_m} |z^* - s|\right)}$$

$$RDFW_i(x, y, z) = \frac{X_m Y_m \beta(x, y, z)}{\int_0^{X_m} \int_0^{Y_m} \int_0^{Z_m} \beta(x, y, z) dx dy dz}$$



$$S_{max,i} = T_{pot} RDFW_i$$

$$\alpha(\psi_m) = \frac{1}{1 + \left(\frac{\psi_m(x, y, z, t)}{\psi_{m,50}}\right)^p} \Rightarrow S_i(\psi_m) = \alpha(\psi_m) S_{max,i}$$

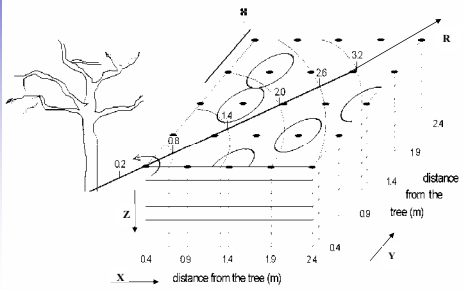
HOW TO FIND MULTI-DIMENSIONAL ROOT WATER UPTAKE PARAMETERS?

- Conduct multi-dimensional experiment
- Construct a multi-dimensional root water uptake model
- Integrate root uptake into multi-dimensional water flow model
- Compare experimental with numerical data
- Minimize their residuals



PARAMETER OPTIMIZATION BY INVERSE MODELING

EXPERIMENTAL LAYOUT OF THREE-DIMENSIONAL SOIL MOISTURE MEASUREMENTS (ALMOND TREE)



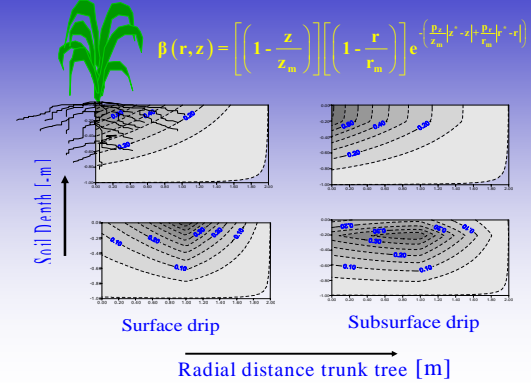
Parameterize root water uptake model

$$\beta(x, y, z) = \left(1 - \frac{x}{X_m}\right) \left(1 - \frac{y}{Y_m}\right) \left(1 - \frac{z}{Z_m}\right) e^{-\left(\frac{p_x}{X_m}|x-x_c| + \frac{p_y}{Y_m}|y-y_c| + \frac{p_z}{Z_m}|z-z_c|\right)}$$

$$S_m(x, y, z) = \frac{X_m Y_m \beta(x, y, z) T_{pot}}{\int_0^{X_m} \int_0^{Y_m} \int_0^{Z_m} \beta(x, y, z) dx dy dz}$$

Unstressed Normalized Root Water Uptake, at a point

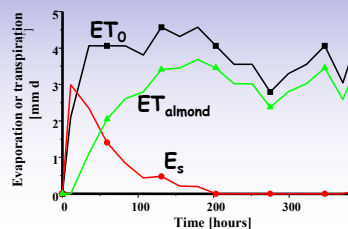
Flexibility of two-dimensional root water uptake model



Computation of T_{pot}

$$T_{pot, almond} = ET_{almond} - E_s$$

$$ET_{almond} = K_c ET_0$$



Water Uptake under Water-Stressed Conditions

Water Stress Response Function

$$\alpha(h_m) = \frac{1}{1 + \left(\frac{h_m(x, y, z, t)}{h_{m,50}}\right)^p}$$

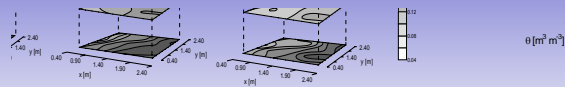
Actual Water Uptake:
 $S(h, x, y, z) = \alpha(h) S_m(x, y, z)$

Actual Plant Transpiration:

$$T_{n,alm} = \frac{1}{X_m Y_m} \int_0^{X_m} \int_0^{Y_m} \int_0^{Z_m} S(h, x, y, z) dx dy dz$$

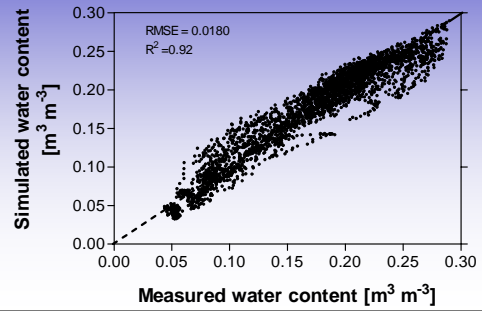
THREE-DIMENSIONAL SOIL MOISTURE OBSERVATIONS

(Vrugt et al., 2001)

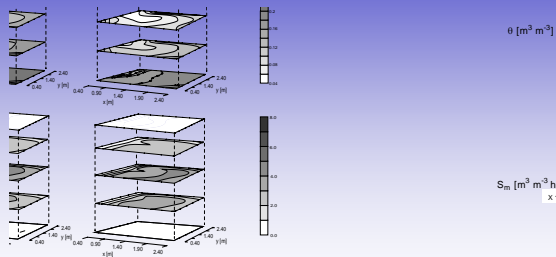


Optimization minimizes residuals of measured and simulated water content values

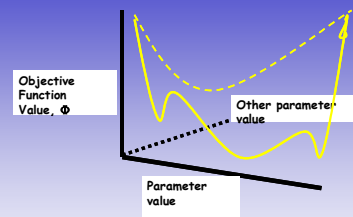
HYDRUS-3D



SIMULATED THREE-DIMENSIONAL SOIL MOISTURE AND ROOT WATER UPTAKE DISTRIBUTION



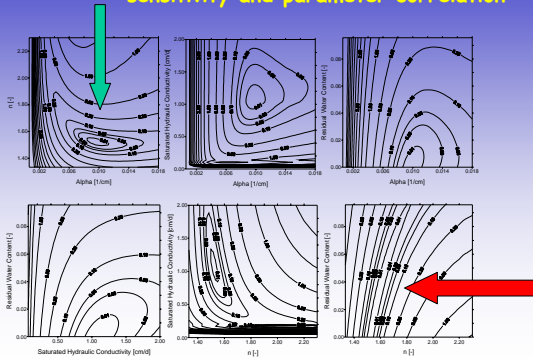
RESPONSE SURFACE ANALYSIS:



- o Objective function contains residuals;
- o Shows local and global minima;
- o Analysis of uniqueness of inverse problem

RESPONSE SURFACE ANALYSIS:

Can also be used to investigate parameter sensitivity and parameter correlation



Well-posed inverse problems:

- Test for global and local minima
- Test for unique solutions
- Independently measure parameters that are not sensitive to solution
- Do not estimate highly correlated parameters
- Include independently-measured information to objective function
- Minimize number of optimized parameters
- Minimize measurement errors
- Estimate model error
- Compare uncertainties of optimized parameters

OTHER APPLICATIONS OF INVERSE MODELING:

- Other soil hydraulic properties techniques, such as evaporation method, suction infiltrometer method and instantaneous profile method;
- Estimation of solute and heat transport properties;
- Estimation of root water and nutrient uptake parameters;
- Effective field soil properties, and in multi-layered systems;
-

LIMITATIONS:

Non-uniqueness
Instability

- Inverse problems are not necessarily well-posed;
- Selection of weighting factors;
- Parameter estimates are valid for experimental range only;
- Method requires a lot of experience

BENEFITS:

- Mandates marriage of experimentation with numerical modeling;
- Method is consistent, I.e. estimated hydraulic functions are used in model predictions;
- Uses transient measurements, as in real world;
- Relatively fast method, and lends itself for automation

